



Bergvesenet

Postboks 3021, N-7441 Trondheim

Rapportarkivet

Bergvesenet rapport nr	Intern Journal nr	Internt arkiv nr	Rapport lokalisering	Gradering
5359				

Kommer fra ..arkiv
Falconbridge

Ekstern rapport nr

Oversendt fra

Fortrolig pga

Fortrolig fra dato:

Tittel

Evje-Iveland prosjektet 1967. Diverse rapporter.

Forfatter

Dato

År

Bedrift (oppdragsgiver og/eller oppdragstaker)

Sulfidmalm AS

Kommune

Fylke

Bergdistrikt

1: 50 000 kartblad

1: 250 000 kartblad

Evje
Iveland

Aust-Agder

Østlandske

Fagområde

Dokument type

Forekomster (forekomst, gruvefelt, undersøkelsesfelt)

Geologi

Råstoffgruppe

Råstofftype

Malm/metall

Sammendrag, innholdsfortegnelse eller innholdsbeskrivelse

A/S Sulfidmalm.

Evie-Iveland project 1967.

Field Report on Map Units E11. E12. D11. D12.

A composite map has been composed of map units E11. E12. D11. and D12.

Various lithological groups have been defined on the map, but the rock types within these groups do not vary from each other a great deal. i.e. unit 1. is predominantly amphibolite. Unit 2. hornblende gneisse and amphibolite and dioritic gneisse, with an area of pegmatites in the north. Unit 3. mainly amphibolite and including the Melland gabbro. Unit 4. amphibolite and hornblende gneisse with a chain of gabbro bodies trending. NE/SW.

TOPOGRAPHY.

About 80-90% of the area is rugged - this usually having a dense forest cover - mainly coniferous. The flat ground is mostly swamp and marsh.

MAIN ROCK TYPES.

- Group 1. Hornblende gneisse, dioritic gneisses and amphibolite.
" 2. Gabbros.
" 3. Pegmatites.

Group 1.

The amphibolite and hornblende gneisse and dioritic gneisses show much the same character when foliated. They can pass into one another both along and across the strike. They are usually fine to medium grained. The amphibolites consist essentially of plagioclase and hornblende.

Common mineral associations in amphibolites are:

Hornblende, plagioclase and quartz.
Hornblende, plagioclase and quartz biotite.
Garnets may or not be present as may ore sulfides.

It seems that they have undergone a complete recrystallization. The amphibolites are cleavable and in many localities it can be seen to split up along wavy surfaces.

In the foliated amphibolites there is a definite segregation of feldspar and hornblende into layers that vary in thickness from less than 1 cm to several cm's - this banding can not be explained by an original bedding and it must be a deformation banding, differential movements causing a mechanical separation of minerals. In certain places one finds that these differential movements have not been able to develop and the rock is massive usually having a granoblastic texture.

In one locality in a massive amphibolite one finds an isolation of clumps of dark constituents in light. It seems that feldspar/quartz substance in the amphibolites undergo a migration - eventually leading to a leaching out of all light constituents over a largish area which becomes enriched in dark constituents often ultrabasic hornblendite. The rock here was not entirely massive but the black constituents had a slight alligement parallel to the general foliation.

The amphibolite, hornblende/dioritic gneiss suits often contains small bands of igneous appearance - these are more or less parallel to the banding and are probably related to the granitization of the area.

The hornblende and dioritic gneiss show most of the characteristics of the foliated amphibolites.

Group 2.

Gabbros occur in all units except the most westerly. The largest is the Melland rock of igneous appearance - within it occur patches of ultrabasics.

Regarding the question as to whether the gabbros are akin to the original basement or related to a later phase of activity and thus escaping to a great extent a deal of metamorphism and recrystallization.

On one side of the block structure of the gabbro and patches of coarser gabbro within the finer gabbro point to a complete recrystallization.

To the south in unit 4 the gabbros were fresh, coarse-grained and with a typically igneous texture. Contacts were difficult to find but one such contact was found at locality (20 Ell). Here the gabbro was medium/coarse grained but as the border was neared the gabbro became finer grained and finally approached a sharp contact with hornblende gneiss. Patches of gneiss had been broken up and incorporated into the gabbro - so here at least the gabbro seems to be related to a later phase. Some of the gabbros to the south of Melland appear to be more acid than the Melland gabbro - perhaps due to 1. a differentiation - i.e. gabbros here at a higher level. 2. Assimilation. 3. A more intermediate intrusion.

Thus this problem of the gabbros is complicated - it seems that some at least of the smaller gabbro bodies are related to a younger period of intrusive activity, and are younger than the Melland gabbro.

Taking the Melland gabbro, a number of observations point to the fact that it is an older type.

1. It is foliated at its edges.
2. Coarser gabbro patches in the finer gabbro suggest complete recrystallization.
3. Block structure suggests recrystallization.
4. Is not fresh in appearance.
5. Dykes that cut the foliation in the hornblende gneisses extent into the gabbro. (These dykes will be discussed later.)

Thus it seems possible that we have two phases of gabbros.

Group 3.

Pegmatites in the area are quite extensive. Most of them were seen to be flat lying. It would seem that the pegmatites in this area are not magmatic but are related to the granitization of the area. There is a lack of feeder channels and the pegmatites have constant border facies - i.e. fine grained quartz/plagioclase muscovite rock, plagioclase graphic granite and large sheets of biotite oriented perpendicular to the border lines.

They are very abundant in amphibolite areas.

IGNEOUS DYKES IN THE BASEMENT.

These were seen in all the map units but were most abundant in E11 around the shores of Kjettevann.

The rocks are fine grained and have a high colour index. They all trend more or less N/S and vary in thickness from a few cm,s to 2-3 mt's.

In the field their mineralogy is much like the amphibolites.

They must have been intruded after the gneisses had been metamorphosed and foliated because they can be seen to be cutting the foliation, but they are earlier than other things i.e. quartz filled cracks parallel to the foliation, tension cracks, local recrystallization and two phases of pegmatites.

1. gneisses
2. dyke
3. quartz cracks parallel to foliation
4. tension cracks
5. local recrystallization
6. pegmatite 1
7. pegmatite 2

MINERALIZATIONS.

Nickel mineralizations were found at two localities within the four map units, at Mølland and at Skripeland.

Smaller areas of mineralization were found at other localities mainly consisting of disseminated sulphides.

At Mølland the ore was confined to the border facies of a gabbro. The chief ore body being in the form of brecciated vein running N/S 1 m thick. D.M.G. tests for nickel proved positive here.

At Skripeland the ore rock is basic and fine grained, thin sections will determine if it is akin to amphibolite or gabbro. To the south this ore rock borders against a fine grained gabbro (diabase) and to the north there seems to be a transition from the ore rock into amphibolites and hornblende gneiss.

GEOLOGISTS IMPRESSION.

As this was the first area mapped the impression one gets of the history of the area is naturally limited until further adjoining areas have been mapped but nevertheless some certain points seem to be clear.

1. Deposition of basement and intrusion of gabbros.
2. Metamorphism and complete recrystallization which introduced foliation into sediments and into borders of gabbro zones.
Regional folding.
3. Intrusion of a series of basic dykes.
4. Gabbro intrusions.
5. Pegmatite introduction.
6. Faulting and jointing.

This may be modified as more evidence becomes available from other areas.

EVJE- IVELAND 1967.

Nickeliferous rocks.

Mølland:- here the rock is a coarse grained massive gabbroic of igneous appearance. It has been described by Vogt (1923) as a hypersthene norite containing the following minerals.

30% Plagioclase An60

35% Hypersthene Fe20

30% Diopside-augite

5% Hornblende

This rock contains 0,09% NiO. The distribution of this rock is unknown it seems to have a patchy occurrence and exhibits insensible transitions into other gabbroid rock types.

Thus a massive type carrying actinolite - like hornblendes also occurs in connection with the ore. On the east side of Kjettevann the gabbroes are usually rather fine grained. Farther south more coarse-grained types are met with. On the west side of Kjettevann the rock carries no pyroxene but it is still massive: the approximate mineral composition is:

40% Plagioclase An40

45% Hornblende

15% Biotite

Ore minerals.

Skipeland, the ore-rock occurs as a small patch inside a granitized gneiss area. The massive, fine grained gabbroid rock exhibits sharp contacts. To the north it borders against a banded gneiss, to the south it borders against what looks like a dike of diabase.

Geologist's Report on Large Square
by

Viggo H. Wiik, July 1967.

ES	FS
E6	F6

General:

The area is moderately exposed. A thin layer of ground morene covers most of the solid rock, but sufficient outcrops are in general available to permit a reliable mapping of the main rock types.

Rock Types, Description and Interpretation:

Basement:

An old, supracrustal basement is recognized, which has served as a host medium for the various younger intrusions and mineralizations. This basement is made up of amphibolite, banded gneisses and darker and lighter dioritic hornblende gneisses. A volcanic origin is suggested for these basement rocks.

In many parts of the area, evidence is found of the basement rocks having suffered deep-seated regional metamorphism leading to partial anatexis and migmatitic phenomena like flow folding and agmatitic structures.

Intrusive Amphibolite:

In some places, e.g. Loc. 22 E6, a dark, finegrained amphibolitic rock is seen to intrude the basement rocks, cutting their foliation and sometimes containing xenoliths of foliated and flow-folded basement gneiss. The intrusive amphibolite is normally structureless, although a foliation may be present. Intrusion of a gabbroic magma must have taken place sometime after the peak of regional metamorphism, and the intruded material later recrystallized to an amphibolitic rock, which cannot by itself be distinguished from the basement amphibolites. So far, only very small dikes and bodies of this intrusive amphibolite have been observed. It is important to keep this fact in mind when reading the geological map, however, because much of what has been mapped as massive, granoblastic amphibolite might well belong to a younger phase of amphibolite, intruded into the basement after peak-metamorphism.

Ore Diorite:

In the western part of the area, rocks occur which look much alike the Ore Diorite at Flåt, described by Barth and Bjørlykke. As is the case at Flåt, specks of sulphides are normal constituents of the ore diorite also in this area. The relation of the ore diorite to the basement rocks is not yet clear. It seems as if the ore diorite was emplaced as fairly large magmatic intrusions at a relatively early stage. In any case, considerable tectonic activity must have post-dated the

Ore Diorite (cont'd):

ore diorite emplacement, as is indicated by fold structures and metamorphic phenomena in ore diorite in some places, such as at Loc. 17 E5. In this road section, the ore diorite appears very inhomogeneous, having light dioritic bands in alternation with elongate ultrabasic bodies. Gradual transition to finegrained, normal ore diorite is seen in the same outcrop. One specimen, 18 E5, taken from within the mapped ore diorite body on map unit E5, is seen in the microscope to be a relatively fresh Hyperstene-Gabbro.

Coarse Hornblende Gabbro:

This rock occurs in a large massif in the western part of E6. The rock is black and white spotted with an average grain size of 5 mm. The main minerals are plagioclase and dark hornblende. To judge from the general appearance, the rock is recrystallized. Directional structures are absent from this rock except in places where local shearing has occurred, frequently in association with pegmatite emplacement. The massive Hornblende Gabbro fractures to form large cubic blocks - an expression of its massive, igneous character. A dike of massive, coarse Hornblende Gabbro, 1-2 ins. wide, is seen to cut foliated ore diorite at Loc. 19 E5. This evidence, together with the non-deformed character of the coarse Hornblende Gabbro in general, indicates a relatively late emplacement of this rock.

"Prestvandite":

A local name has been given to this rock in order to distinguish it as a rock type. The "Prestvandite" is a finegrained, light brownish to greenish grey rock. The fresh variety is structureless, but a planar structure is easily developed, and it sometimes gets a characteristic appearance with streaks of biotite in a massive, grayish groundmass. This rock has been found as a sizeable body only at Loc. 25-26 F5. Here it intrudes migmatitic basement gneiss but is itself brecciated by granite and granite pegmatite. Elsewhere, "Prestvandite" has been found as dikes, characteristically related to fault zones. ^{Under} the microscope, "Prestvandite" has a granoblastic texture. The major constituents are plagioclase, quartz, microcline and biotite.

Dolerite:

In a few places, small, mostly dike-like intrusions have been observed, of a dense, greenish black massive rock, which ^{under} the microscope shows relict hyperitic texture with laths of plagioclase in an amphibole-rich, granoblastic matrix. These rocks seem to represent a fairly young episode of gabbroic intrusion. They brecciate the coarse Hbl-Gabbro at Loc. 9 E6, and they show no signs of having experienced later tectonization.

Granitic Rocks:

Whereas granite pegmatites are ubiquitous in the entire area, other granitic rocks are met with only in the northern part. Here a variety of granitic rocks occur, from porphyroblastic augen gneiss through finegrained, sugary granitic rocks with or without gneiss structure, to medium-coarse grained magmatic granite. Much work remains if the interrelationships among the granitic rocks are to be sorted out. What is clear at present is that the granitic rocks came into the area at a relatively late stage, and that there are several stages in the emplacement history of the granitic rocks.

Pegmatites occur as vertical dikes and as flat-lying, cake-shaped bodies often with a gentle northerly dip. No mineralogical differences can be found between the two types of occurrence.

Granitization of basement rocks is seen to occur along certain shear zones, such as at 9^XF6. Within the areas of granitic gneiss, relict zones of dioritic and amphibolitic material are sometimes found, and this is taken as an indication of the metasomatic origin of at least some of the granitic gneisses.

Mineralizations:

Old prospects have been found at the localities 21, 22, 16 and 12 F6 and 22 E5. In addition, notable sulphide contents are observed in rocks from the localities 19 E5 and 8 E5. No distinct pattern emerges from the distribution of these mineralizations within the present Large Square, nor has a common source been discovered. The mineralized zones of map unit F6 must be investigated more closely in Phase 2 of the Iveland-Evje Project.

Molybdenite occurs in minor amounts in the granitic gneiss at 16 F5 and as larger crystals in the pegmatite north-west of 28 F5.

Structure:

The structural relationships seem to be very complicated, and so far, no significant discoveries have been made. The basement rocks probably varied with regard to primary structures, the massive amphibolite having none, whereas the banded gneiss had a primary stratification. The various dioritic hornblende-gneisses probably developed their foliation during early metamorphism. All these structures were thoroughly re-modelled during the stages of ultrametamorphism. On the basis of the structures found in some occurrences of what is thought to be ore diorite, it is suggested that this rock also experienced some kind of regional dynamo-metamorphism. It must remain an open question as to whether the ore diorite was emplaced prior to the migmatization period, or whether there was a later phase of regional metamorphism.

Structure (cont'd):

As indicated on the map, the area is dissected by numerous fault zones, some of which have cataclastic rocks associated with them. It is interesting to note that in many cases, the foliation is seen to change locally towards a position parallel to the fault.

27th July, 1967.

VW/pb

Jan 2.1.1968
EPRC

Geologists Report on Large Square

E3	F3
E4	F4

G4	H4

by
V. H. Wink, August -67

The area is very rugged, with big faults cutting the young granites as well as the old basement rocks. The granite areas are well exposed, but this is generally not the case with the basement gneiss areas.

To the west, medium-grained granitic gneiss is the dominant rock type. Only in a few places have minor areas of amphibolite been observed to break the homogeneity of the granitic rock. The granitic gneiss may be characterized as homogeneous on a macroscopic scale, whereas on a mesoscopic scale pinpoint flow fold structures may still be seen, leaving the impression of this rock having formed through granitization of migmatitic basement-gneisses.

In the southwestern part, a wide zone of augengneiss occurs, probably forming a more or less continuous zone transitional between the westerly granitic gneiss and the zone of dioritic hornblende-gneiss running along the

central N-S axis of the large square. Minor bodies of intrusive granite occur within the basement gneiss zone, and to the east, massive coarse to medium-grained granite becomes the dominant rock type. This granite is clearly intrusive, carrying numerous minor xenoliths of ~~amphibolite~~ basement rocks, mostly amphibolite. Also larger units of amphibolite are situated as islands in this sea of igneous granite.

It is interesting to note that xenolithic bodies of the rock type previously named "Prestvandite" occur in the granite in the eastern part of the large square.

The granite dominates also map unit E4, but here continuous zones of amphibolite start to occur between ridges of granite and in H4, there is a dominance of basement rocks. A prominent fault runs through H4, separating a suite of migmatitic amphibolite and hornblende gneiss from a wide area of sheared and granitized gneisses to the east. Photogeological features particularly to the north of H4 indicate a considerable displacement along this fault.

Pegmatites are not as numerous as further south in the Inland-Eye area, but they are not rare. Contrasting Barth's statement pegmatites have been observed intruding the granite, thus at least one phase of pegmatite emplacement must post-date ^{the} granite intrusion.

Mineralizations.

Important sulphide mineralization has been found only in one place within the large square. This is at loc. 15E4, where two ^{prospecting pits?} small shafts have been worked. Here pyrrhotite, chalcopyrite and pyrite occur as disseminations in a dense to medium-grained greyish rock much like the rock elsewhere called ore diorite. Outside the mineralized zone only amphibolites and dioritic hornblende gneisses are seen.

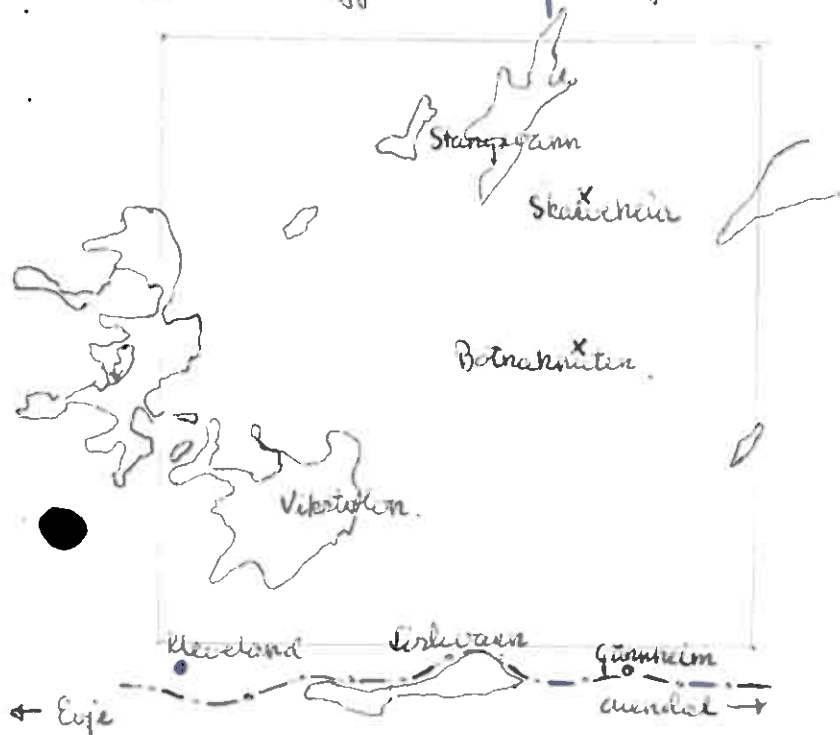
Structure:

It is interesting to note that in east, on H4, the foliation strikes NNE with a W-ely dip, whereas in the gneiss-granite area in the W, on E3 and E4 there is a more or less similar strike of the foliation, but here the dips are to the E. Does this mean that the granite has intruded the central zone of a syncline?

Seen 1978.
30.12.67

Rapport omfattende map units: G5. G6. H5. H6.

Området ligger nord for Ege-Aundalsen, fra Kleivland til Gjønnheim



Se skisse. Terranget stiger bratt opp fra Aundalsen.

Området er også jevnt over meget kupert. Dette gjelder særlig H4. H5 & H6, hele den sørlige del av G6 og det nordvestre hjørnet av G5.

Rundt 10% av området er vann, noe liknende gjelder myrene. Området ellers er skogbevokst, løvskogen dominerer

noe. Den kvartære overdekning varierer. I øst mangler den flere steder og ellers er det for det meste et tynt dekke av bunnmorene.

Regional geologi

De dominerende bergarter er granitt, granittisk gneiss, pyegneiss, amfibolitt, diorittisk gneiss og diorittiske hornblendegneisser. Dessuten har en granittpegmatitter og granittiske ganger.

Amfibolitten er fin til middels kornet. Den har en granoblastisk teksture og er som regel massiv. Bortsett fra steder er foliasjon utviklet. Mineralsammensetningen er amfibol, plagioklas ± kvarts ± biotitt. Dette komplekset med amfibolitt dekker den NW del av området (se kart).

Den diorittiske gneissen som en har i H.U. G5 ~~viser~~ stor likhet med amfibolitten. Den er massiv, finkornet, granoblastisk og er noe mer rik på biotitt. Det siste kan skyldes granittmassivet i NW.

Hornblende-gneissen som begrenser amfibolitten mot ESE, ligger under amfibolitten og strekker N til NNE. Hornblendegneissen er finkornet og av mere diorittisk sammensetning. Den er godt foliert. Den er biotittrik og på sine steder noe granittisert.

Granittisk gneiss utgjør størstedelen av området, rundt 60%, og begrenser hornbl. gneissen østover. Den er massiv og forholdsvis homogen over hele området, selv om en kan finne hornblendegneiss irste her og der. Disse er da meget biotittrike. Den granittiske gneissen varierer en del i kornstørrelse, fra fin til over middels. Der hvor en har pegmatitter kan en f.eks. ikke skikkelig se grensen mot gneissen.

I den sørøstlige delen av H.U. H6 har jeg øyegneiss-utvikling flere steder.

Intusius-granitt har jeg et par steder i området. I det nordvestre hjørnet er det en middels til grovkornet granitt, noen steder har den blitt opp og inneholder fragmenter av amfibolitt. Mere finkornet granitt har jeg i det sørøstlige hjørnet av H6. Og så denne inneholder fragmenter av amfibolitt. Deruten er det en mindre kupp av granitt i det NW hjørnet av H5. Fin til middels kornet.

Pegmatittene er av granittisk sammensetning og av den enkle typen. De har en skarp avgrensning til rikebegarten når de opptrer i amfibolitten og hornblendegneissen, men ikke i den granittiske gneissen. De granittiske gangene må antas

å høre sammen med granittintrusjonene.

Amfibolitten, diorittisk gneiss og hornbl.gneissene må høre til basement. Granittintrusjonene er betydelig yngre. Hvor det granittiske gneisskomplekset skal plasseres er jeg ikke riktig sikker på, sannsynligvis sammen med basementbergarten.

Strukturer

Feliasjonen er ganske konstant over hele området; ligger rundt 220-260.

En markert forkastningslinje går gjennom hele området, fra St. Stangerann, vest for Vikstølerrann og til vestenden av Lislebann. Andre markerte "faults" grupperer seg

a) i strekkretningen

b) nesten øst-vestlig retning. Kan ha forbindelse med Ege-Gårnheim fault.

Rapport på "large square": 69.610.149.110

Det blir referert til map units m.

Området ligger rundt Vassvatn, ca 2 km nord og sør for veien Fikstad - Vigjårdal.

Terranget er småkupert, for det meste dekket av skog. Rundt 7% av arealet består av vann, 5% er myr og bare rundt 1% er dyrket. Et mer eller mindre tynt lag av brennmorene dekker det meste av området, imidlertid er det som regel nok bløtninger av fast fjell tilgjengelig.

Regional geologi

Hardbergene i området er båndgneiss, hornblendegneiss, amfibolitter, hornblendegabbro og grunnfiske bergarter.

Båndgneissene er finkornede, har vekslende lag av diorittisk og amfibolittisk sammensetning, og er på sine steder meget sterkt migmatisert (eks. lokalitet 149, det for bro over Vassvatn). De lyse og mørke lag har skarpe grenser. Noen steder er båndene meget smale, andre steder kan en ha ganske mektige soner med f.eks. diorittisk s. setn. Gneissen kan være biotittrik.

Amfibolitten er fin til middels kornet og flere steder med partiell anatexis. Det samme gjelder hornblendegneissene.

På mappeint "G10" og den sørvestre del av G9
har en opptreden av grovkornede hornblendegabbroer.
Her har en også amfibolittar av skiftende utbredelse.
Jeg har ikke fått særlig mye ut av dette
området (bare 1 tårn), så disse gabbroene og deres
opptreden børde kan ende kartlegges nærmere.

De granittiske bergartene er av flere typer. Intusiv-
granitt har jeg i H10. Den er middels- til grovkornet
og har noen steder bruddstykker av amfibolitt.

I samme område opptreer også en finkornet granitt
(39 H10). Samme bergart har jeg også i G9 (lok. 36 G9).

I H9 opptreer også en fin til middels kornet granitt.
Jeg har inntrykk av at den har tenget opp parallelt
foliasjonen (conformt), slik at den ligger som
relativt (tykke ~~f~~) tynne flak over gnissbergartene. Den
har og så bruddstykker av denne i seg.

Andre bergarter som opptreer i området er ultra-
basitter, en noe kisholdig intusiv-gangbergart (muligens
or-diorite), intusiv amfibolitt, doleritt, granitt-
pegmatitt, granittiske ganger (apittar) og dissesten
lyse ganger av mer kvartsdiorittisk sammensetning.

Av de ultrabasiske bergartene var en særlig ~~måke~~
seg lokalitet 27 G9. Bergarten en finner der, på vestsiden
av Vassvatn, kan muligens være det Bath har kalt
Bahiaitt. Den inneholder sulfider av Fe, Cu og Ni.
Dimetylglucosimtest var positiv.

Den finkornede kisholdige b. art (or-diorite²) opptreer
vid lok. 24 G9. Dette er i en "shear-zone".

Merk finkornet amfibolitt som må ha vært intrusiv
og observert ved to lokaliteter:

10 H10 Viser en noe over 1/2 m mktig gjennomskjærende
gang

669 Viser amfibolitt som har intrusivt båndgneiss

Granitpegmatitt påtreffes over hele området, med
skarpe grenser til de omgivende bergarter. De er for
største delen av den enkle typen. En finkornet aplitt,
av granittisk s. setn., som skjærer gjennom granit-
pegmatitt er observert ved lok B H10. Denne må
således være yngre.

Et sattrakk ved dette området, i motsetning til
det første området mitt i nord, er forekomsten av
gjennomsettlende ganger av kvartsdiorittisk sammen-
settning. Disse har ikke alltid skarpe grenser, det
virker som løsningsene har vært aggressive og naget
noe med rikkbergarten. En stau her kanskje overfor
et mobilisat (metasom²) som er tilført fra områder
med delvis oppsmeltning (meget høy metamorfosegrad)

Dette kan passe godt med de sterkt migmatiserte
gneisser og amfibolitter en påtreffer i området.
En må her hatt en ganske høy grad av metamorfose
her.

Strukturer

Et par kraftige "tectonic lines" går over området
i retning N 60 til N 80. Den ene går gjennom G10 og H10,
den andre gjennom det skråttlige området av H10. I
tilknytning til den siste, er en "shearzone" observert.

Map units C9.C10.D10

1. Mikrorittene børde kartlegges nærmere, særlig for å forstå deres oppførsel
2. Jeg har inntrykk av at de ultrabasiske b. antene opptrer langs forkastnings soner.

Området er ganske rikt på disse ultrabasiske
Ved å kartlegge disse bedre, kunne en kanskje
få litt system i hvordan de opptrer, om de
opptrer som ganger eller mindre legemer,
og hvor i området de opptrer hyppigst.